

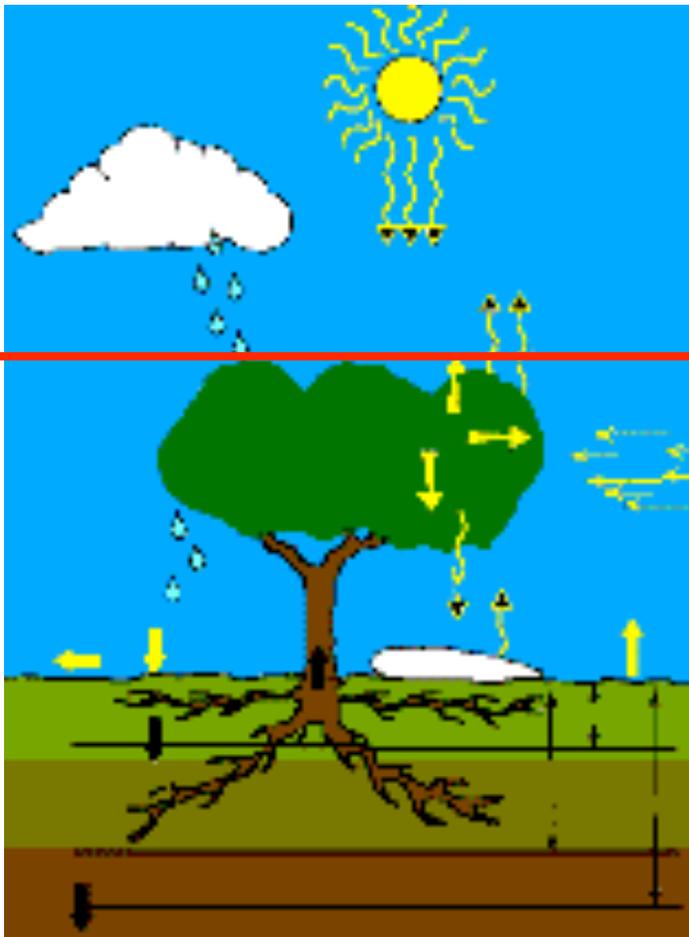


Capitalizing on 1D RT models to interpret remote sensing data from 3D structurally heterogeneous vegetation systems

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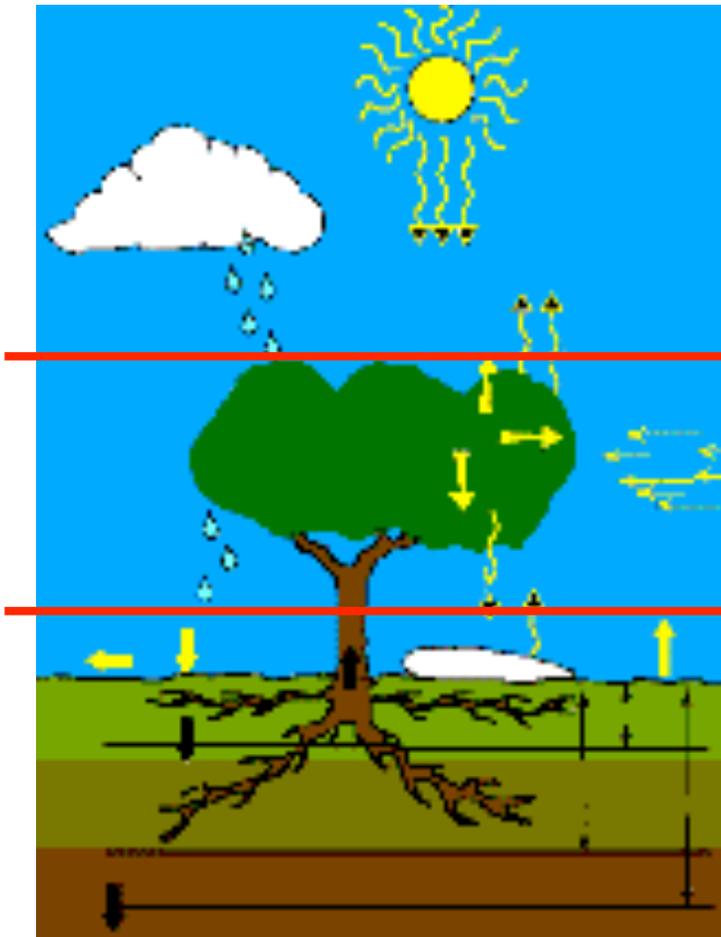
How does radiation redistribute energy between the atmosphere and the biosphere?



- The “surface” corresponds to the boundary condition of RT atmospheric problem

Need to understand and represent the albedo of that “surface”

How does radiation redistribute energy between the atmosphere and the biosphere?



- The “surface” corresponds to the boundary condition of RT atmospheric problem

Need to understand and represent the albedo of that “surface”

- The “surface” corresponds to the upper boundary condition of the vegetation plus soil RT and other problems

Need to understand and represent the RT processes yielding the distribution of energy below that “surface”, e.g., transmitted fluxes

3-D structural effects and short term climate: the snow case with ECMWF/NCEP

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glossary on off

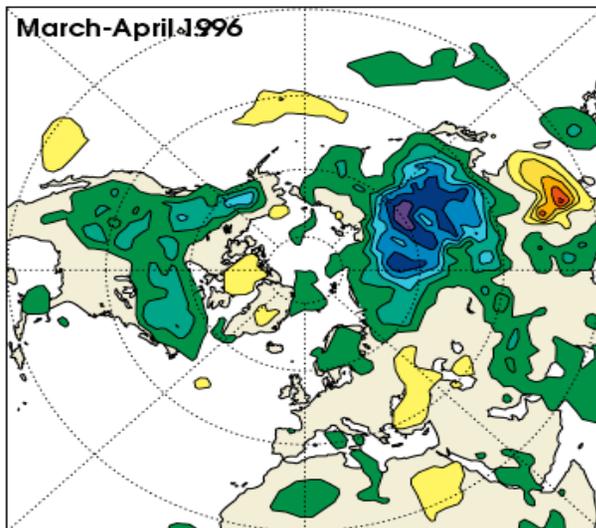


FEATURES

Ref: Viterbo and Betts, 1999, JGR

Everyone Complains About the Weather...

Betts and his **BOREAS** colleagues observed that, in the spring, daily **weather** forecasts significantly underestimated air **temperatures** over the **boreal** forest, sometimes by as much as 10—15°C (18—27°F) (Viterbo and Betts, 1999). Additionally, the BOREAS team found that predictions of cloud cover over the boreal region were often far off the mark. Everyone complains about the weather, but how could the forecasts be so wrong so often?



The scientists noticed a pattern that confirmed their earlier suspicions: the temperature forecasts were farthest off in late spring when snow was on the ground and grew more accurate after the snow melted. From summer through fall, the weather **models** matched actual measurements more

1 ◀ ▶ 3

This map shows the average errors in the European Centre for Medium-Range Weather Forecasts at 850mb (roughly equivalent to an altitude of 1500m) for March and April of 1996. The predictions, made five days in advance, were compared to actual measurements. The 1996 model did not include the adjustments to forest albedo.

(Figure from Viterbo, P. and A.K. Betts, 1999: The impact on ECMWF forecasts of changes to the albedo of the boreal forests in the presence of snow. J. Geophys. Res. (in press, BOREAS special issue). Courtesy A.K. Betts)



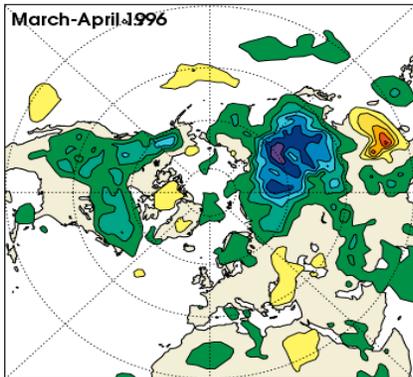
“...weather forecasts significantly underestimated air temperatures over boreal, sometimes by as much as 10-15 C...”

Ref: <http://eobglossary.gsfc.nasa.gov/>

3-D structural effects and short term climate: the snow case with ECMWF/NCEP

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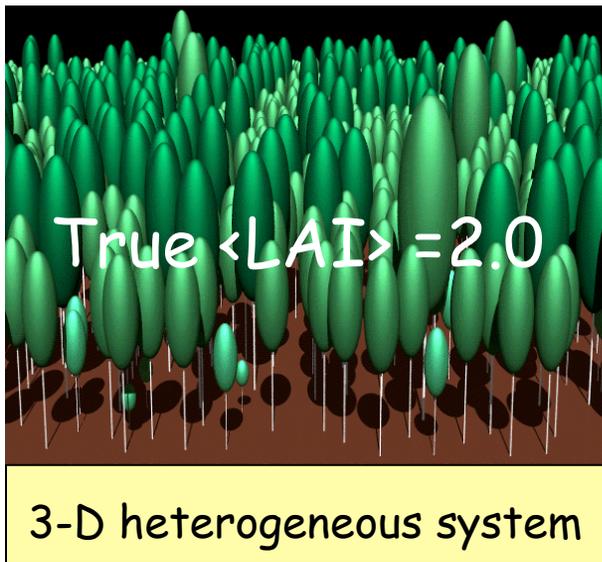
“...—the BOREAS team found that the models were overestimating albedo (the amount of light reflected by the surface). ...”



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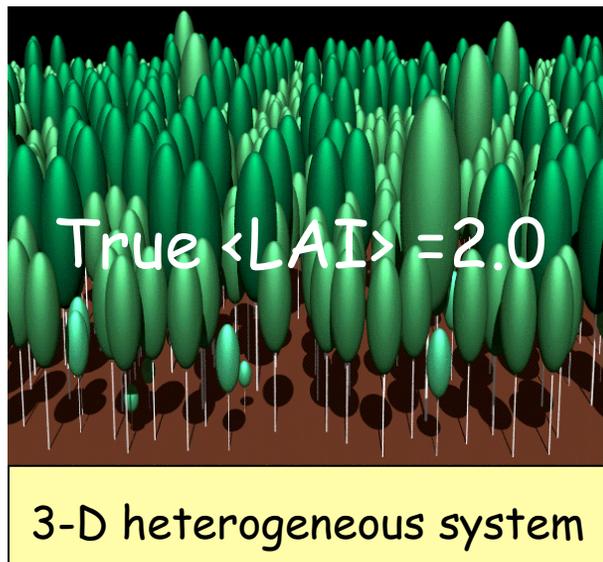
Using “true”, i.e., domain-averaged, optical depth and other true radiation transfer (RT) state variables ($\langle X \rangle$) in a 1D RT scheme can only yield seriously erroneous radiant flux estimates



Direct transmission at 30 degrees Sun zenith angle,

$$T_{1-D}^{direct}(\langle LAI \rangle) = \exp\left(-\frac{\langle LAI \rangle}{2\mu_0}\right) = 0.312$$

Using “true”, i.e., domain-averaged, optical depth and other true radiation transfer (RT) state variables ($\langle X \rangle$) in a 1-D RT scheme can only yield seriously erroneous radiant flux estimates



Absorption in the visible (PAR) domain

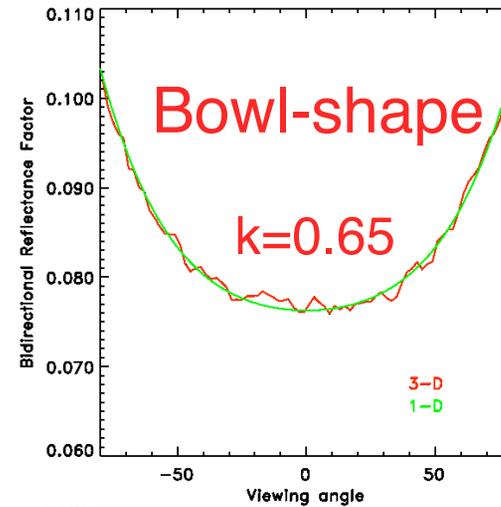
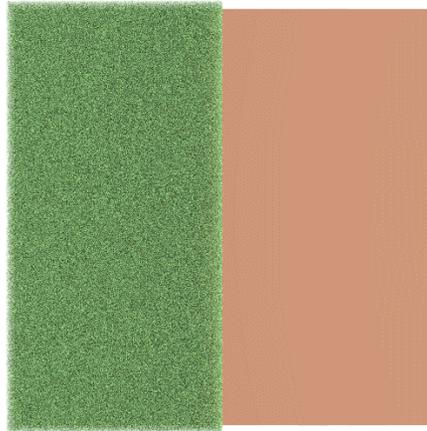
← 1-D 2-stream model

← 3-D MC model

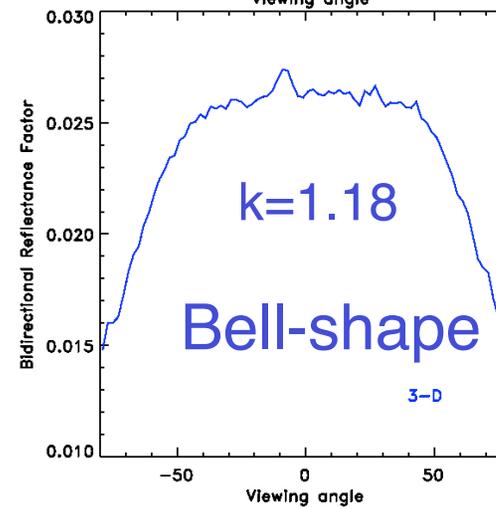
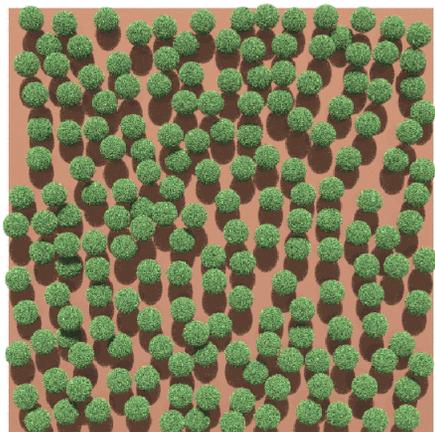
Albedo (DHR) in the near-infrared domain



Angular signatures in the red region



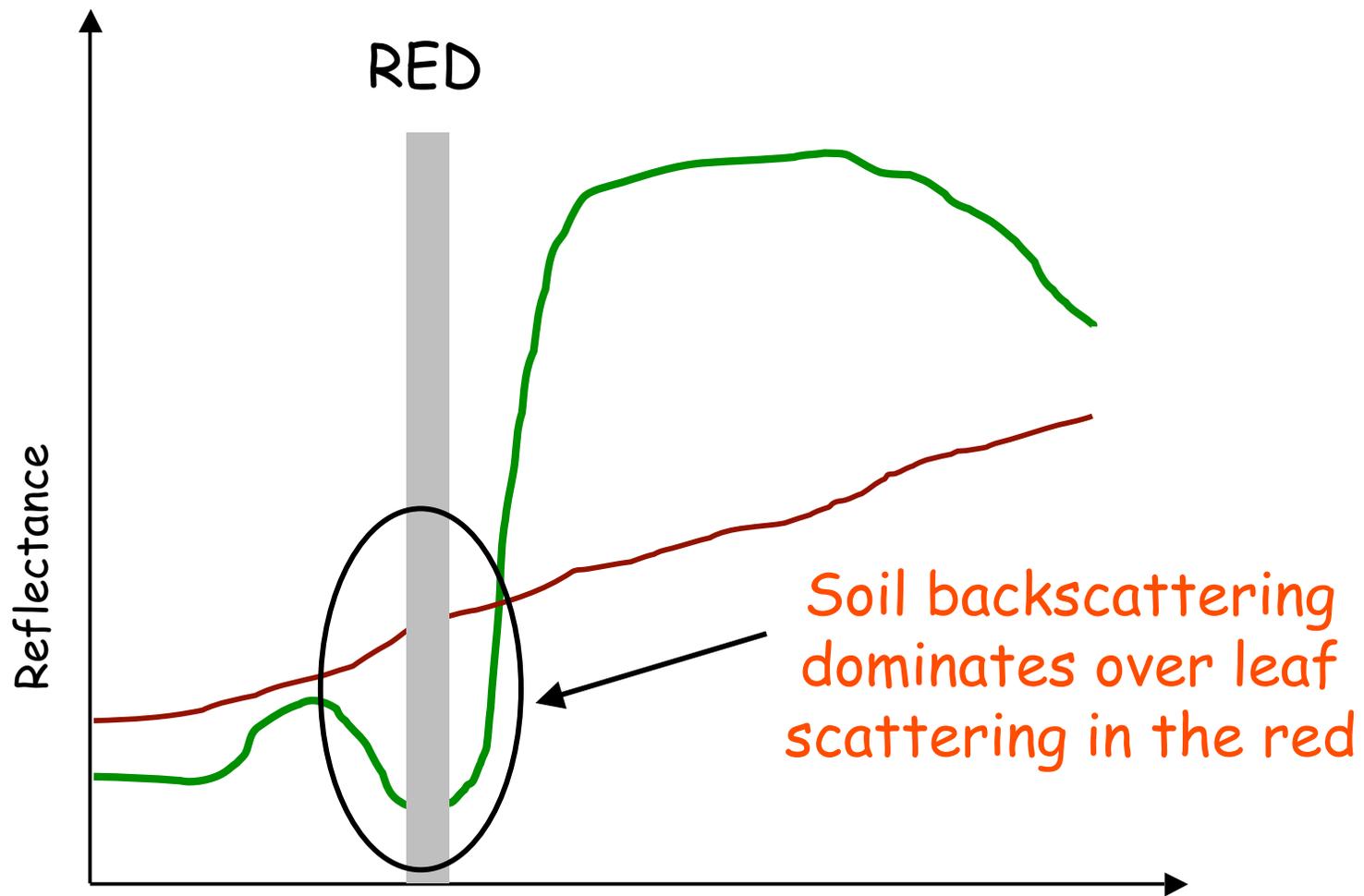
1-D



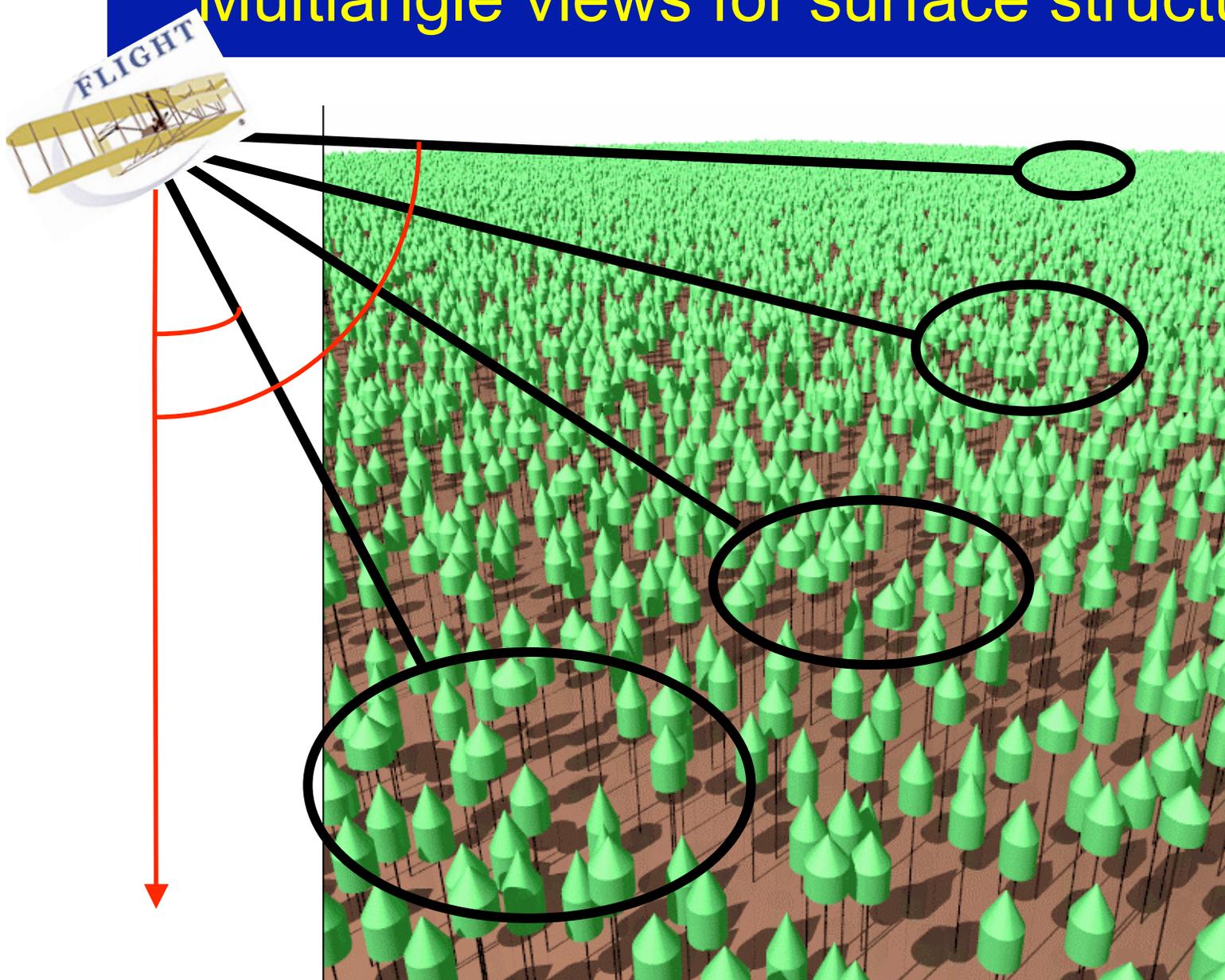
3-D

Ref: Pinty et al. (2002) IEEE TGRS

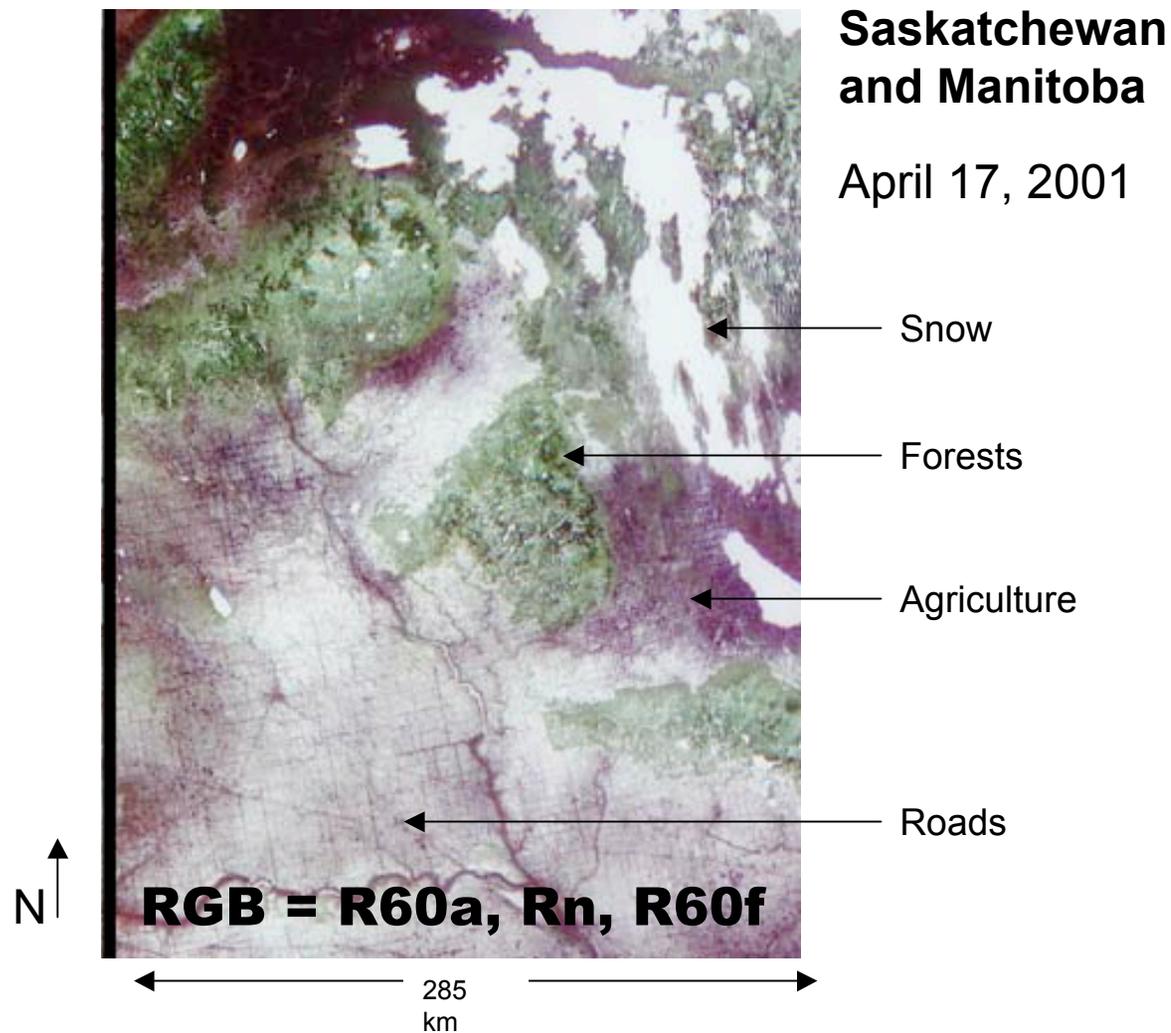
Spectral contrast between the vegetation canopy and the background



Multiangle views for surface structure



Surface structure from MISR/Terra



Ref: http://www-misr.jpl.nasa.gov/gallery/galhistory/2001_may_30.html

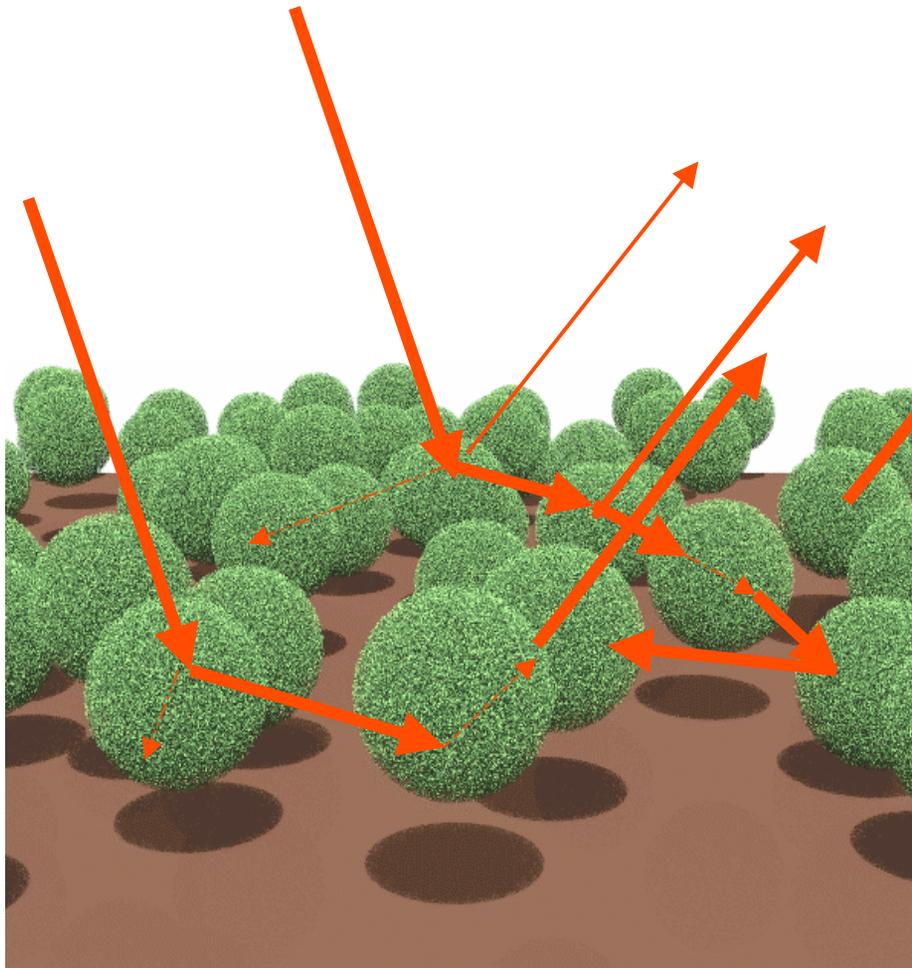
Main objective

- Retrieve information on state variables from 3-D vegetation systems such as forests while preserving the integrity of the structural effects.

Approach

- Optimize the use of 3-D MC simulations for the most significant structure-induced effects
- Couple with 1-D equivalent RT solutions for the less contributing factors

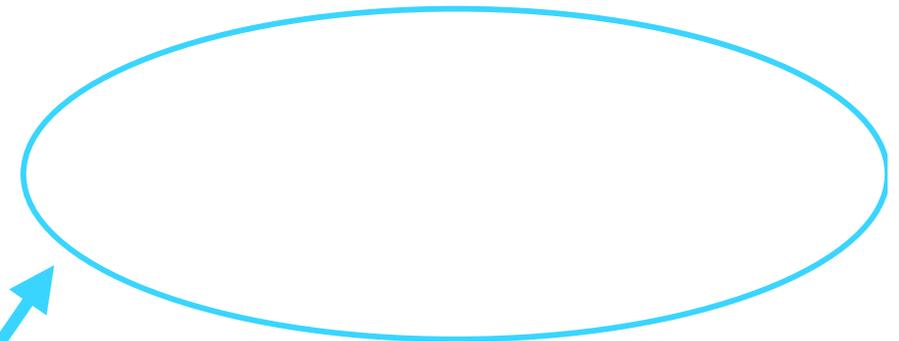
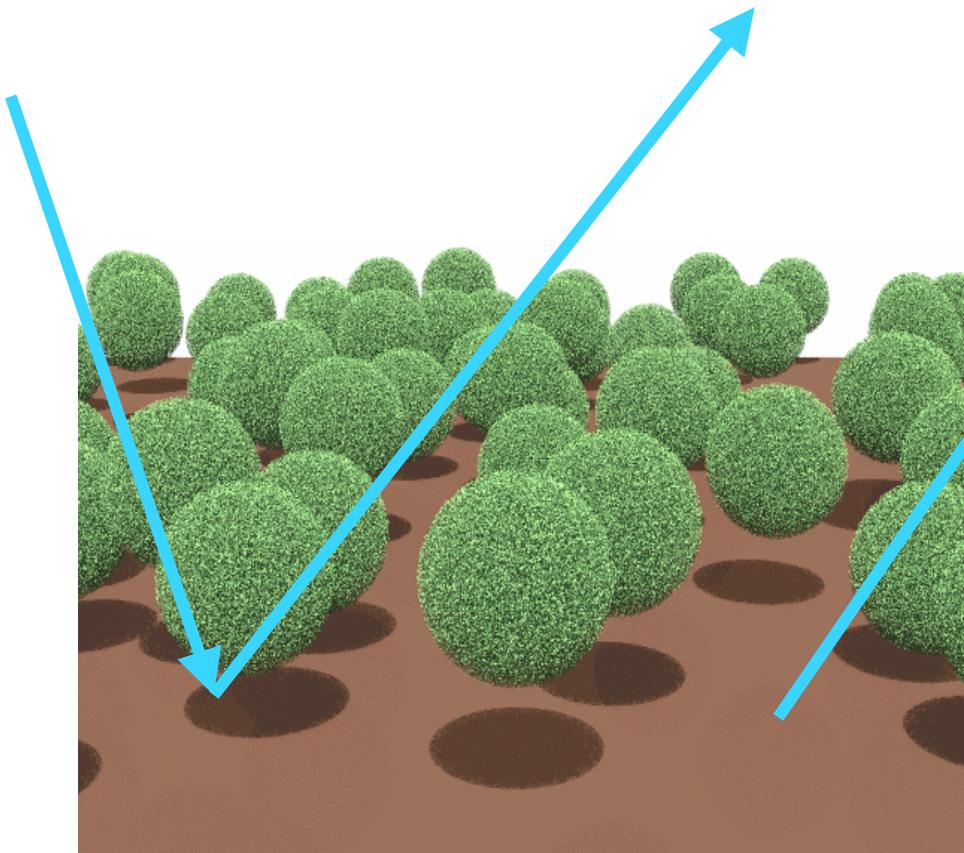
Decompose the complex problem into simpler problems to solve



Black Background contribution

- 1) It regulates the *absorption* processes associated to vegetation photosynthesis
- 2) Strongly depending on the *density* of green vegetation

Decompose the complex problem into simpler problems to solve

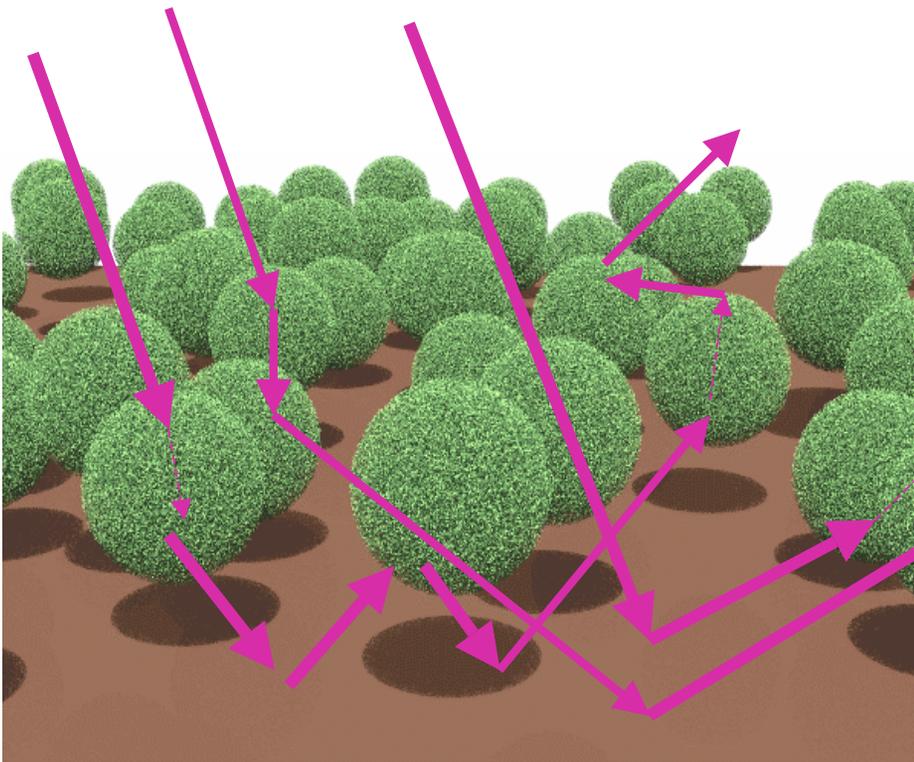


Black Canopy contribution

1) No absorption process by vegetation associated with this wavelength independent contribution

2) Strongly controlled by *3-D distribution* of vegetation extinction coefficient

Decompose the complex problem into simpler problems to solve



1) Controlled by multiple scattering events between the background and the canopy

2) Mostly **negligible** in the visible domain of the solar spectrum

Snow cover

Near-Infrared

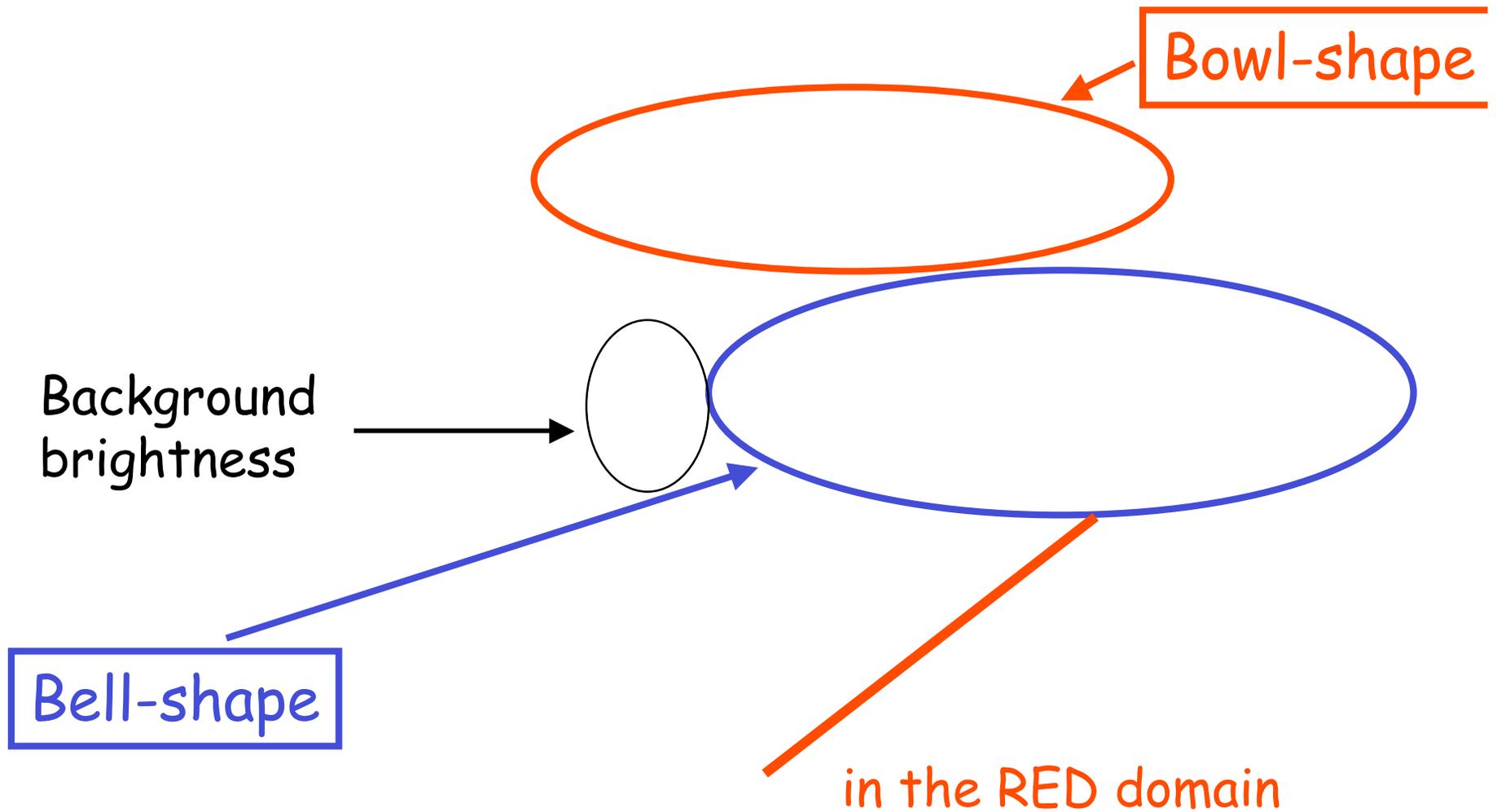
Red

Black
Background
contribution

Black Canopy
contribution

Multiple Collided
contribution with
background

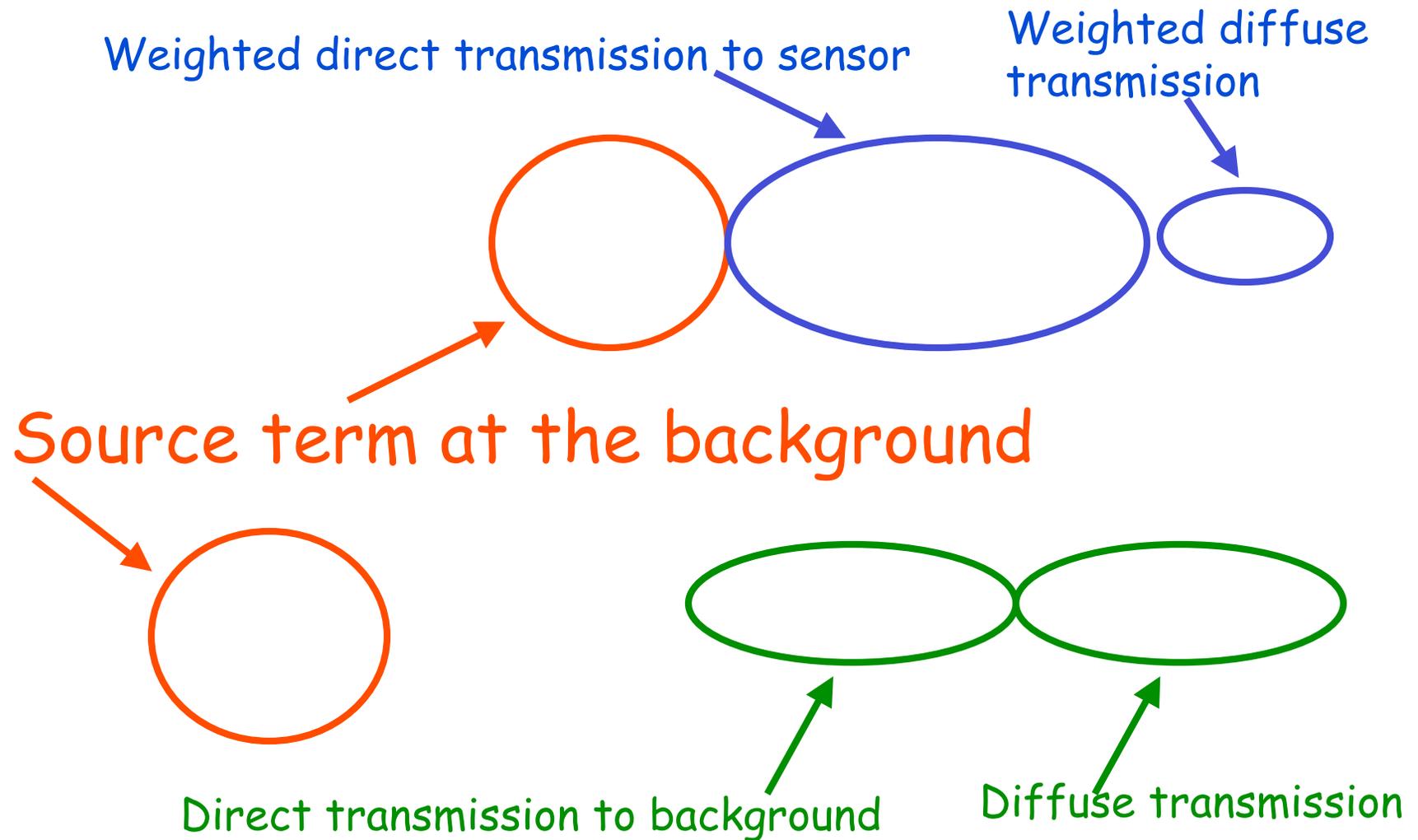
Decompose the complex problem into simpler problems to solve



Use a 1-D solution for the background-canopy multiple scattering contribution

1. Use azimuthal average of the diffuse radiance fields scattered by the background
2. Assume a homogeneous plane-parallel turbid medium based representation

Parameterization of the equivalent turbid medium layer: All scattering orders >1 with the background



Parameterization of the equivalent turbid medium layer: First order of scattering with the background

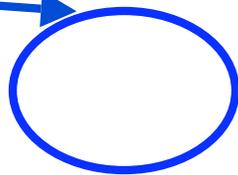
$$\rho_{bgd1}^{\parallel Coll}(z_{toc}, -\mu, \mu_0) \approx \exp(-\widetilde{LAI}/2 |\mu|) f_0(-\mu, \mu_0)$$

$$+ \exp(-\widetilde{LAI}/2 \mu_0) g_0(-\mu, \mu_0)$$

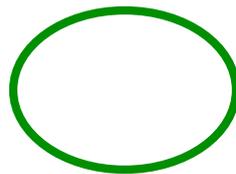
$$+ h_0(-\mu, \mu_0)$$

Direct
transmissions
with effective
LAI value

Azimuthal average of
diffuse transmission



Azimuthal average of
background BRF shape



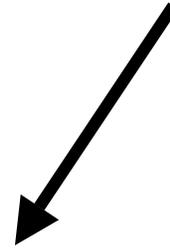
Requirements from a 1-D turbid medium

- 3 (effective) state variables:
 1. *Optical depth: LAI*
 2. *single scattering albedo :*
Leaf reflectance+ Leaf transmittance
 3. *asymmetry of the phase function*
Leaf reflectance/transmittance
- 2 boundary conditions:
 1. *Top: Downward flux from the atmosphere*
 2. *Bottom : Upward flux from the background*

Estimates of the effective radiative quantities

\tilde{LAI}

is retrieved from inversion of the exponential attenuation law against Black Canopy (wavelength-independent) contribution from 3-D simulations



\tilde{LAI}

\tilde{LAI}

is lower than the true domain-averaged

Estimates of the effective radiative quantities

$\tilde{}$ and $\tilde{}$
and

are retrieved from inversion of the R and T_{diff} fluxes associated with the Black Background contribution from 3-D simulations and given \tilde{LAI}



$$\sum_{\mu_0} |\overline{T}_{bgd1}^{Coll}(z_{bgd}, \mu_0; \mathbf{B}, \mathbf{O}, \mathbf{S}, LAI_{scene})|$$

Reflected flux

$$- \overline{T}_{bgd1}^{||Coll}(z_{bgd}, \mu_0; \tilde{r}_l, \tilde{t}_l, \tilde{LAI}(\mu_0))|$$

Diffuse flux

Ratio of effective/true domain-averaged values

Variable	Sparse	Medium	Dense
\tilde{LAI}	0.35	0.45	0.76
$\tilde{\quad} + \tilde{\quad}_{(NIR)}$	0.82	0.84	0.88
$\tilde{\quad} / \tilde{\quad}_{(NIR)}$	4.65	5.76	7.14

Final simplified expression for the radiance fields from any complex land surface

$$\begin{aligned}
 \rho_{coupled}^{total}(z_{toc}, \Omega, \Omega_0) \approx & \rho_{veg}^{Coll}(z_{toc}, \Omega, \Omega_0; \mathbf{B}, \mathbf{O}, \mathbf{S}, LAI_{scene}) \\
 & + \rho_0 [\rho_{bgd}^{UnColl}(z_{toc}, \Omega, \Omega_0; \mathbf{B}, \mathbf{S}, LAI_{scene}, k, \Theta) \\
 & + \rho_{bgd}^{Coll}(z_{toc}, -\mu, \mu_0; \mathbf{B}, \tilde{\mathbf{O}}, \tilde{LAI}_{scene}, k, \Theta, \rho_0)]
 \end{aligned}$$

Monte Carlo solutions (points to ρ_{veg}^{Coll})
 Black Background (points to ρ_{veg}^{Coll})
 Black Canopy (points to ρ_{bgd}^{UnColl})
 Analytical solution (points to ρ_{bgd}^{Coll})
 Multiply scattered with background (points to ρ_{bgd}^{Coll})

Performance of the analytical solution

Snow-covered background

Typical background

Sun angle at 30°

Performance of the analytical solution

Snow-covered background

Typical background

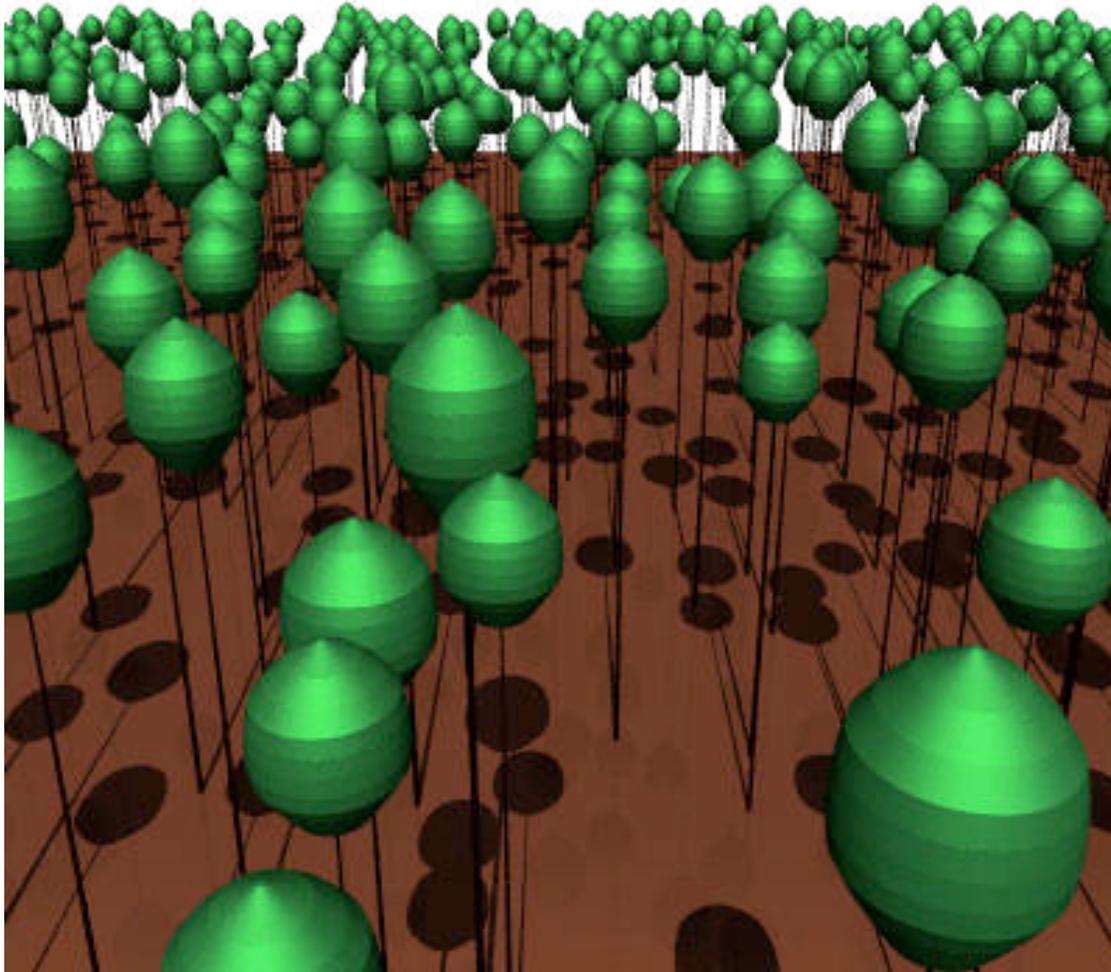
Sun angle at 60°

Performance of the analytical solution

Inversion strategy

- Use ecological knowledge, e.g., allometric equations to design a large set of canopy scenarios once for all:
take advantage of RAMI results

Examples of geophysical scenarios



Inversion strategy

- Use ecological knowledge, e.g., allometric equations to design a large set of canopy scenarios once for all: **take advantage of RAMI results**
- Estimate the Black Background and Black Canopy contributions associated with all scenarios: **use 3-D MC simulations**
- Estimate the effective state variable values of the 1-D problem : **LAI from the Black Background and r_l and t_l from the Black Canopy contributions**
- Adopt an inversion scheme minimizing the distance between measurements and simulations: **Adjoint model of the 1-D representation**

What did we gain?

- Inversion scheme based on limited MC simulations: **Black Background** and **Black Canopy** components only are requested
- The background brightness value is solved as a continuous variable during the inversion
- Allows us adding constraints on the **Black Canopy** contribution as a function of time
- Allows detection of fast changes in the background conditions, e.g., snow and snow melting

